



Reconstructed monthly river flows for Irish catchments 1766–2016

Paul O'Connor¹ | Conor Murphy¹ | Tom Matthews² | Robert L. Wilby²

¹Irish Climate Analysis and Research Units,
Department of Geography, Maynooth
University, Maynooth, Co. Kildare, Ireland

²Department of Geography and
Environment, Loughborough University,
Loughborough, UK

Correspondence

Paul O'Connor, Irish Climate Analysis and
Research Units, Maynooth University, Rm
1.9., Laraghbryan House, County Kildare,
Ireland.

Email: pkoconnor@gmail.com

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Abstract

A 250-year (1766–2016) archive of reconstructed river flows is presented for 51 catchments across Ireland. By leveraging meteorological data rescue efforts with gridded precipitation and temperature reconstructions, we develop monthly river flow reconstructions using the GR2M hydrological model and an Artificial Neural Network. Uncertainties in reconstructed flows associated with hydrological model structure and parameters are quantified. Reconstructions are evaluated by comparison with those derived from quality assured long-term precipitation series for the period 1850–2000. Assessment of the reconstruction performance across all 51 catchments using metrics of MAE (9.3 mm/month; 13.3%), RMSE (12.6 mm/month; 18.0%) and mean bias (−1.16 mm/month; −1.7%), indicates good skill. Notable years with highest/lowest annual mean flows across all catchments were 1877/1855. Winter 2015/16 had the highest seasonal mean flows and summer 1826 the lowest, whereas autumn 1933 had notable low flows across most catchments. The reconstructed database will enable assessment of catchment specific responses to varying climatic conditions and extremes on annual, seasonal and monthly timescales.

KEY WORDS

hydrological modelling, Ireland, reconstruction, river flow, time series

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1 | INTRODUCTION

Continuous, long-term river flow records are needed for evaluations of hydro-climatic variability and change, historical extremes and catchment processes (Machiwal and Jha, 2006). They also underpin water management and provide a means of stress-testing existing and planned systems to a range of variability and past droughts (Wilby and Murphy, 2019). Unfortunately, there are few continuous and homogeneous river flow records spanning a century or more (Mediero *et al.*, 2015). Instead, available records are often impacted by confounding factors or large amounts of missing data (Wilby *et al.*, 2017).

Various techniques exist for extending observations by reconstructing river flows. This typically involves forcing statistical or conceptual hydrological models with long-term precipitation and temperature/evapotranspiration data provided by reanalysis (e.g. Kuentz *et al.*, 2013; Brigode *et al.*, 2016) or long-term historical data sets (e.g. Jones, 1984; Spraggs *et al.*, 2015; Crooks and Kay, 2015; Rudd *et al.*, 2017; Hanel *et al.*, 2018; Smith *et al.*, 2019; Noone and Murphy, 2020). Others have leveraged international data rescue initiatives to generate gridded historical weather variables (Casty *et al.*, 2007). Whilst these kinds of information have been used to reconstruct river flows in parts of Europe (e.g. Moravec *et al.*, 2019), they have yet to be deployed in the British-Irish Isles.

Here, we develop a data set of reconstructed monthly river flows for 51 catchments across the island of Ireland back to 1766. This was achieved using gridded historical meteorological data, bias corrected to contemporary observations in each catchment. These data provided the input to a conceptual hydrological model and an artificial neural network (ANN), both of which were trained and verified using river flow observations. In addition, we use recently rescued precipitation data to evaluate model reconstructions for selected catchments during the period 1850–2010. The following sections describe the catchments, data sets and modelling approaches, before we present the derived reconstructions.

2 | DATA PRODUCTION METHODS

2.1 | Catchments and data

Reconstructions were generated for 51 catchments (Table 1 and Figure 1) that are relatively free from artificial influences (following criteria applied by Murphy *et al.* (2013): they have at least 25 years of record and acceptable quality rating curves). The catchments are broadly representative of hydro-climatological conditions across the island, with a recognized under-representation of upland catchments along coastal

margins (Broderick *et al.*, 2019). Urban extent averages <2% of the combined area of all catchments, which individually vary in size between 10 and 2,418 km². However, given the extent of arterial drainage works undertaken in Ireland, it is unavoidable that some catchments have been impacted by such activities. We note which catchments are known to be affected by arterial drainage in Table 1.

Daily flow series were obtained from the Office of Public Works (OPW; <http://waterlevel.ie/>) and the Environmental Protection Agency (<http://www.epa.ie/hydronet/>) and then aggregated to monthly mean flows. The average amount of missing data was <6% across the 51 catchments, with a notable outlier of 31% being the Blackwater at Duarrigle (ID: 18050). Of the total missing days (11% overall), the majority have been previously infilled using rainfall–runoff modelling techniques (Murphy *et al.*, 2013). As the remaining missing data only represented 1% of the total, they were not repopulated.

We use gridded (1 × 1 km) monthly precipitation and temperature series (Walsh, 2012) area-averaged for each catchment, alongside concurrent river flow records, to calibrate the hydrological models (see below). Monthly potential evapotranspiration (PET) was estimated from air temperature and radiation following the method of Oudin *et al.* (2005). We favoured this over more physically based methods (e.g. Penman–Monteith), because the latter have greater data requirements (e.g. wind speed, humidity) that cannot be met over the full duration of the reconstruction period. Instead, the sensitivity of monthly river flow simulations to PET estimation methods was tested for periods with complete variable sets. Six PET estimation methods (Penman–Monteith Penman (1948), Monteith (1965), Blaney and Criddle (1950), Hamon (1961), Oudin *et al.* (2005), Thornthwaite (1948) and Kharrufa (1985)) were evaluated using the hydrological model GR2M. This revealed that the Oudin method performed similarly to the Penman–Monteith method, with an average RMSE of 3.6 mm between flows generated from the two methodologies for five catchments for the period 1974–2000 (equating to 4.5% of mean annual flows).

2.2 | Historical gridded precipitation and temperature data

Casty *et al.* (2007) (henceforth Casty data) produced gridded (0.5° × 0.5°) monthly temperature and precipitation series for Europe covering the period 1766–2000 using non-linear principle component regression of a spatial network of available station data against reanalysis data, with independent predictors used for different variables (Casty *et al.*, 2007). Monthly mean temperature and total precipitation were extracted and averaged for grids overlying each catchment for the years 1766–2000. Quantile mapping (Maraun, 2016) was used to

TABLE 1 Details of the 51 catchments for which flow reconstructions were generated

River ID	Flow station		Waterbody		Arterial		Area		Calibration		GR2M validation scores			ANN validation scores			Ensemble validation scores		
	Name		Name		Drainage		km ²	Years	Years	NSE	KGE	PBIAS%	NSE	KGE	PBIAS%	NSE	KGE	PBIAS%	
3051	Faulkland		Blackwater (Mon)		Yes		143	1976–2000	1976–2000	0.79	0.69	-16.40	0.78	0.67	-14.80	0.80	0.69	-16.00	
6013	Charleville		Dee		Yes		309	1976–2000	1976–2000	0.78	0.75	-12.50	0.82	0.70	-9.10	0.83	0.72	-11.20	
6014	Tallanstown		Glyde		Yes		270	1976–2000	1976–2000	0.77	0.75	-8.00	0.82	0.73	-4.50	0.81	0.73	-7.70	
6030	Ballygoly		Big		No		10	1975–2000	1975–2000	0.86	0.86	-2.10	0.83	0.78	-1.00	0.86	0.83	-1.80	
7009	Navan Weir		Boyne		Yes		1658	1977–2000	1977–2000	0.77	0.73	-9.90	0.81	0.73	-6.40	0.81	0.71	-9.60	
7012	Slane Castle		Boyne		Yes		2,408	1961–2000	1961–2000	0.79	0.74	-11.90	0.83	0.71	-8.60	0.84	0.72	-10.90	
12001	Scarawalsh		Slaney		No		1,031	1961–2000	1961–2000	0.76	0.73	-15.10	0.80	0.81	-11.90	0.78	0.74	-14.50	
14007	Derrybrock		Stradbally		No		115	1980–2000	1980–2000	0.83	0.77	-10.10	0.88	0.79	-8.80	0.86	0.75	-10.30	
14019	Levistown		Barrow		No		1697	1961–2000	1961–2000	0.81	0.74	-12.40	0.86	0.79	-8.00	0.84	0.74	-11.60	
15001	Annamult		Kings		No		445	1972–2000	1972–2000	0.87	0.90	-1.00	0.88	0.82	1.10	0.89	0.87	-0.30	
15003	Dinin Bridge		Dinin		No		140	1972–2000	1972–2000	0.87	0.86	-7.40	0.84	0.75	-8.90	0.88	0.82	-7.80	
15005	Durrow Ft. Br.		Erkina		No		379	1972–2000	1972–2000	0.78	0.68	-11.90	0.83	0.72	-8.50	0.82	0.68	-10.60	
15006	Brownsbarn		Nore		No		2,418	1972–2000	1972–2000	0.87	0.88	-3.90	0.90	0.87	0.30	0.91	0.87	-2.10	
15007	Kilbricken		Nore		No		340	1982–2000	1982–2000	0.82	0.71	-11.30	0.81	0.68	-8.60	0.82	0.70	-10.70	
16008	New Bridge		Suir		No		1,090	1961–2000	1961–2000	0.85	0.84	-6.90	0.86	0.87	-2.40	0.88	0.84	-5.40	
16009	Caher Park		Suir		No		1583	1962–2000	1962–2000	0.86	0.87	-7.50	0.90	0.93	-3.20	0.89	0.89	-6.20	
16010	Anner		Anner		No		437	1973–2000	1973–2000	0.80	0.89	-1.20	0.84	0.92	2.90	0.85	0.91	0.60	
16011	Clonmel		Suir		No		2,144	1962–2000	1962–2000	0.88	0.89	-0.10	0.89	0.94	3.20	0.90	0.90	1.10	
16012	Tar Bridge		Tar		No		230	1969–2000	1969–2000	0.83	0.83	0.80	0.83	0.89	2.00	0.85	0.85	1.30	
16013	Fourmilewater		Nire		No		94	1973–2000	1973–2000	0.69	0.85	-0.80	0.82	0.84	-1.20	0.80	0.86	-1.50	
18002	Ballyduff		Blackwater		No		2,334	1972–2000	1972–2000	0.90	0.91	6.00	0.87	0.86	10.30	0.91	0.92	7.40	
18003	Killavullen		Blackwater		No		1,257	1972–2000	1972–2000	0.91	0.88	0.80	0.92	0.95	3.80	0.93	0.91	1.60	
18006	Cset Mallow		Blackwater		No		1,052	1978–2000	1978–2000	0.88	0.83	-0.20	0.89	0.88	3.80	0.90	0.85	1.10	
18050	Duarrigle		Blackwater		No		250	1982–2000	1982–2000	0.89	0.86	-5.50	0.89	0.84	-5.00	0.90	0.85	-5.30	
19001	Ballea		Owenboy		No		103	1973–2000	1973–2000	0.85	0.76	-16.30	0.82	0.74	-12.60	0.85	0.75	-14.50	
21002	Coomhola		Coomhola		No		65	1976–2000	1976–2000	0.90	0.83	-6.40	0.92	0.89	-4.70	0.93	0.87	-5.30	
22006	Flesk		Flesk (Laune)		No		329	1990–2000	1990–2000	0.84	0.75	-7.20	0.91	0.84	-6.40	0.89	0.80	-6.60	
22035	Laune Bridge		Laune		Yes		560	1992–2000	1992–2000	0.81	0.73	-6.10	0.87	0.84	-6.30	0.86	0.78	-6.40	
23002	Listowel		Feale		Yes		647	1975–2000	1975–2000	0.93	0.90	2.80	0.92	0.87	3.40	0.94	0.89	3.40	
24008	Castleroberts		Maigue		Yes		806	1977–2000	1977–2000	0.89	0.83	-0.40	0.88	0.83	3.30	0.90	0.82	0.90	

(Continues)

TABLE 1 (Continued)

River ID	Flow station		Waterbody		Arterial		Area		Calibration		GR2M validation scores			ANN validation scores			Ensemble validation scores		
	Name		Name		Drainage		km ²		Years		NSE	KGE	PBIAS%	NSE	KGE	PBIAS%	NSE	KGE	PBIAS%
24030	Danganbeg		Deel		Yes		259		1981–2000		0.92	0.90	1.70	0.91	0.90	4.20	0.93	0.90	2.80
25001	Annacotty		Mulkear		Yes		648		1973–2000		0.88	0.82	-1.50	0.88	0.82	0.00	0.89	0.82	-1.00
25002	Barrington Br.		Newport (Mun)		Yes		230		1961–2000		0.91	0.92	-0.90	0.91	0.94	0.70	0.92	0.93	-0.20
25006	Ferbane		Brosna		No		1,163		1961–2000		0.83	0.79	-10.40	0.86	0.80	-6.80	0.87	0.78	-9.50
25030	Scarriff		Graney		Yes		279		1973–2000		0.86	0.81	-6.80	0.86	0.83	-4.10	0.87	0.82	-5.50
25034	Rochfort		L. Ennell Trib		Yes		11		1976–2000		0.81	0.84	-7.90	0.85	0.81	-7.60	0.86	0.82	-8.30
26021	Ballymahon		Inny		No		1,099		1973–2000		0.83	0.86	-4.10	0.84	0.88	-0.60	0.88	0.85	-3.40
26029	Dowra		Shannon		Yes		117		1976–2000		0.86	0.85	-8.80	0.83	0.81	-8.00	0.86	0.84	-8.20
26058	Ballyrink Br.		Inny Upper		Yes		60		1982–2000		0.75	0.82	4.50	0.85	0.82	4.40	0.85	0.82	2.60
27002	Ballycorey		Fergus		Yes		511		1961–2000		0.83	0.74	-8.70	0.84	0.81	-6.90	0.85	0.75	-8.50
30007	Ballygaddy		Clare		No		470		1975–2000		0.89	0.85	-7.50	0.91	0.90	-4.30	0.91	0.86	-6.40
32012	Newport Weir		Newport		No		146		1982–2000		0.90	0.85	-6.00	0.90	0.89	-3.80	0.91	0.87	-4.80
33001	Glenamoy		Glenamoy		Yes		76		1978–2000		0.93	0.95	-0.50	0.87	0.94	0.00	0.93	0.96	0.30
34001	Rahans		Moy		No		1975		1970–2000		0.90	0.92	-4.60	0.89	0.94	-1.90	0.92	0.92	-4.00
35002	Billa Bridge		Owenbeg		No		81		1972–2000		0.86	0.88	4.70	0.87	0.92	5.60	0.88	0.90	5.30
35005	Ballysadare		Ballysadare		No		640		1961–2000		0.89	0.90	-2.30	0.90	0.90	-2.10	0.91	0.90	-2.40
36015	Anlore		Finn		No		153		1973–2000		0.84	0.76	-9.50	0.79	0.65	-10.40	0.84	0.73	-9.80
36019	Belturbet		Erne		No		1,492		1961–2000		0.83	0.85	-5.60	0.79	0.77	-7.60	0.86	0.80	-6.90
38001	Clonconwal		Ownea		No		111		1973–2000		0.93	0.87	-4.00	0.93	0.88	-4.50	0.94	0.88	-3.60
39006	Lennan		Claragh		No		245		1977–2000		0.85	0.88	8.00	0.85	0.87	7.90	0.86	0.88	8.00
39009	Aghawoney		Fern O/L		Yes		207		1973–2000		0.91	0.90	-1.00	0.91	0.88	-2.20	0.92	0.90	-1.20

Note: Included are calibration periods for each catchment, together with logNSE, KGE and PBIAS scores for the validation period (2001–2016) for ANN, GR2M and Ensemble median simulations.

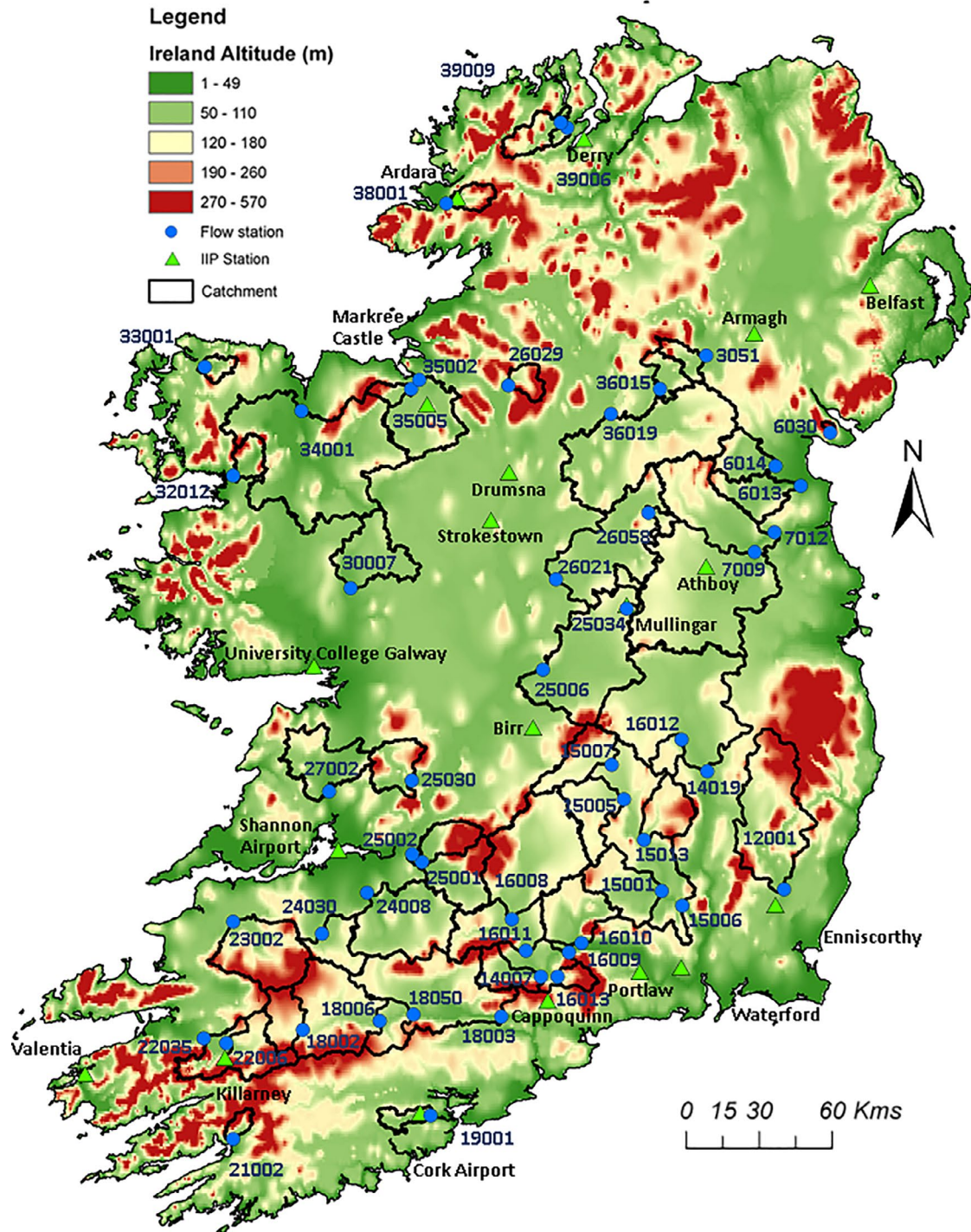


FIGURE 1 The 51 study catchments along with the locations of corresponding flow stations and island of Ireland precipitation (IIP) series synoptic stations

bias correct Casty data to catchment averages using the aforementioned gridded (1×1 km) monthly precipitation and temperature series. We perform quantile mapping by interpolating the empirical quantiles using local linear least square regression to robustly estimate the values of the quantile–quantile relationship between the Casty and observed data for each catchment. For values outside the historical range, a constant correction—equivalent to the highest quantile in that

series—was applied (Boé *et al.*, 2007). Bias correction was carried out on a monthly basis using the ‘qmap’ R package (Gudmundsson, 2016). Sample bias correction plots for nine catchments are shown in Figure 2 (temperature) and Figure 3 (precipitation). Across the 51 catchments, the bias adjustment produced minimal change in mean annual temperature values (-0.15°C). Precipitation corrections were more substantial, with a mean increase of 94.2 mm/year (7.7% of mean annual

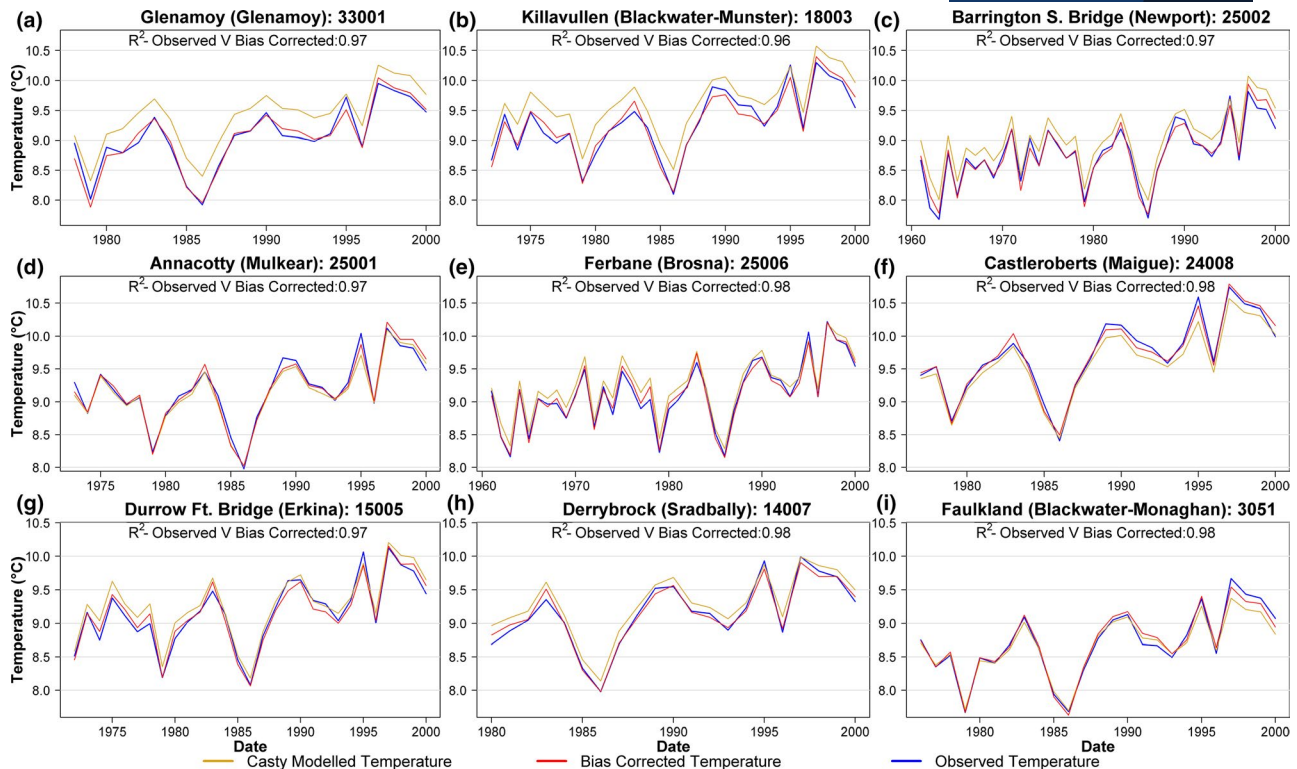


FIGURE 2 Annual bias corrected Casty temperature for nine catchments from the start of the respective observations up until the year 2000. R^2 scores between bias corrected and observed temperature values are also provided

precipitation). Once bias corrected, observed temperature and precipitation were appended to each catchment series to bring values up to 2016. The Oudin method was then used to derive PET estimates from the Casty temperature data for each catchment.

2.3 | Hydrological models and calibration procedures

To ascertain the contribution to uncertainty generated by model structure, two model types were implemented—a conceptual hydrological model (GR2M) and an empirical based Artificial Neural Network (ANN). These models are explained below.

2.3.1 | The GR2M conceptual model

GR2M is a simple water balance model (Mouelhi *et al.*, 2006), originally developed for French catchments, now available via the airGR R hydrological modelling package (Coron *et al.*, 2017). The monthly flow model contains two reservoirs representing a soil store and routing reservoir (Figure 4) governed by two parameters: the production store capacity and groundwater exchange coefficient. GR2M has been widely deployed across diverse catchment types and

applications (e.g. Louvet *et al.*, 2016), including for flow reconstructions (Dieppois *et al.*, 2016).

For each catchment, GR2M was calibrated and validated on observed data before using the bias corrected Casty data to reconstruct flows. A split record for calibration/validation was applied as this allows direct comparison between GR2M and ANN model outputs on a catchment-by-catchment basis. Calibration for all catchments (including a 1-year warm-up period) was undertaken from the start of the flow record up to December 2000. This time interval captures periods of large flow variability ranging from the drought rich 1970s to the flood rich 1980s. Validation was undertaken using the 15 years postcalibration (2001–2016) for all catchments (see Table 1).

Uncertainty in GR2M model parameters was sampled using Monte Carlo methods. For each parameter, 20,000 values were randomly drawn from a uniform distribution of [0–2500] for the production store capacity and [0–2] for the groundwater exchange coefficient. Each parameter set was used to simulate flows for the calibration period (yielding a 20,000-member ensemble). The performance of parameter sets was evaluated using two objective functions to ensure robust performance across the flow regime: the Nash Sutcliffe Efficiency (NSE) (Nash and Sutcliffe, 1970) derived from log transformed flows (logNSE) and the modified Kling Gupta Efficiency (KGE) derived from raw flows (Gupta *et al.*, 2009; Kling *et al.*, 2012). Two steps were then

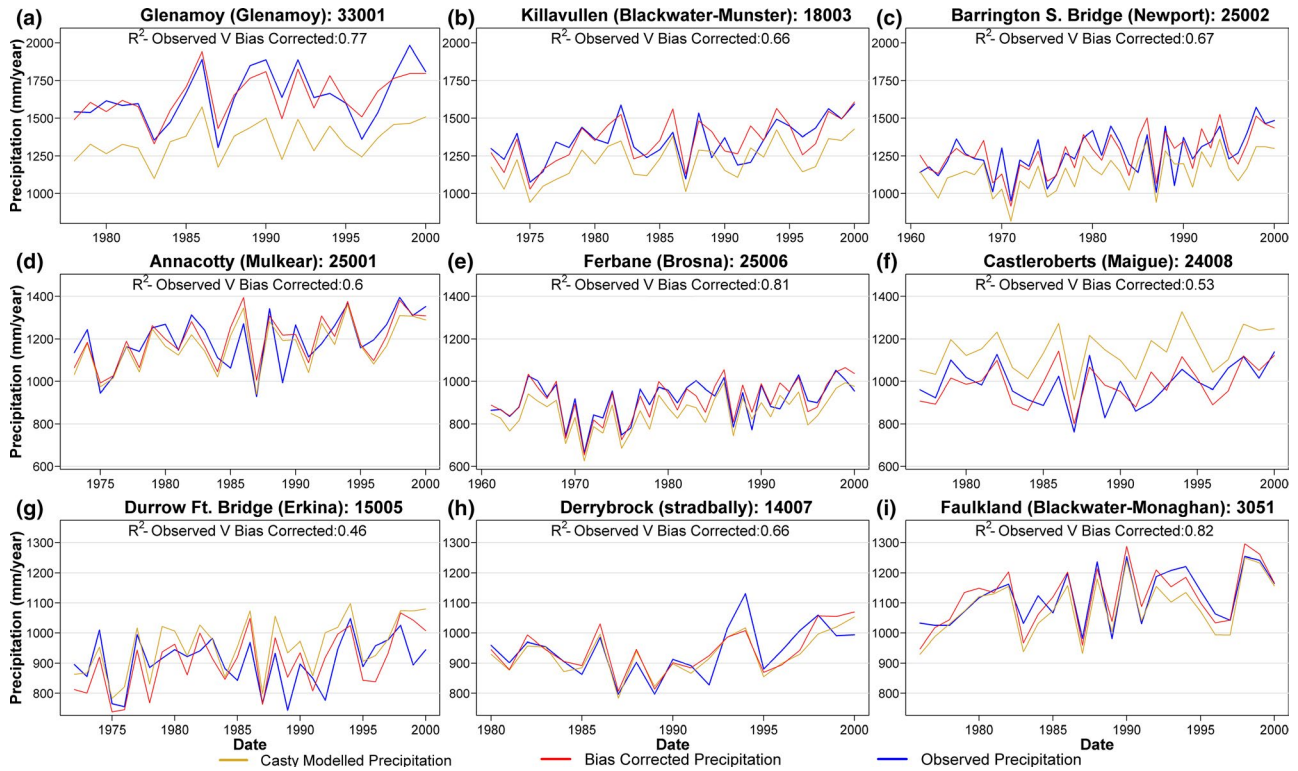


FIGURE 3 Annual bias corrected Casty precipitation values for nine catchments from the start of the respective observations up until the year 2000. R^2 scores between bias corrected and observed precipitation values are also provided

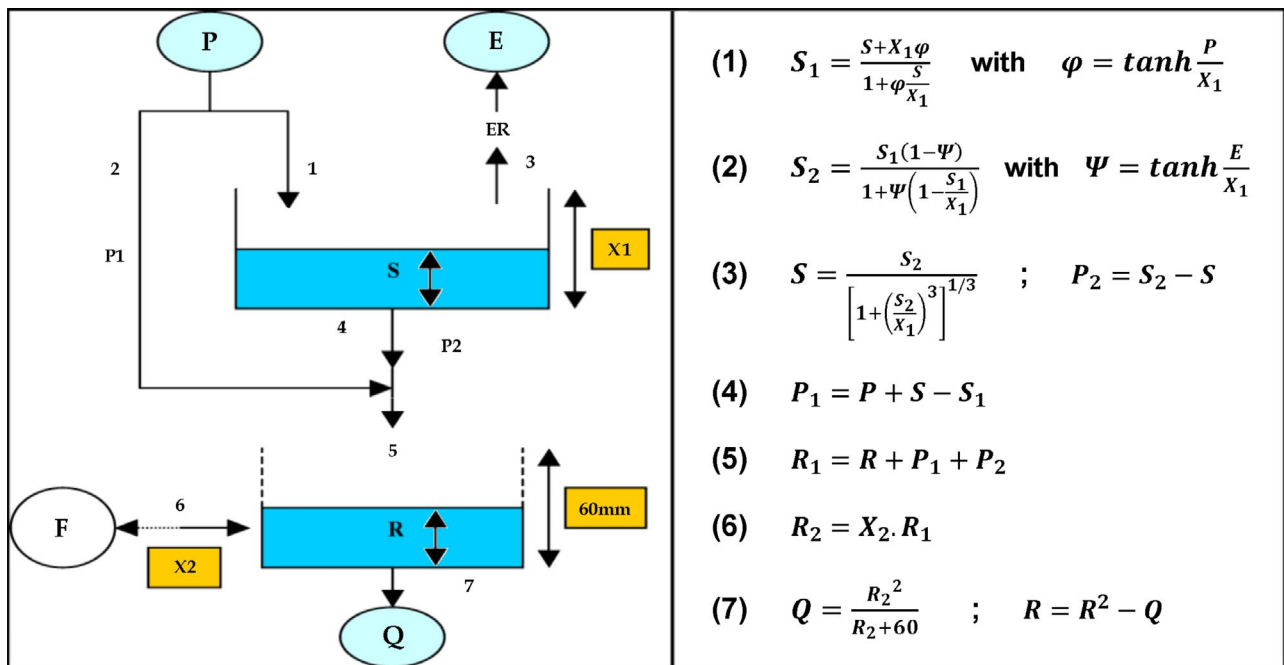


FIGURE 4 Outline of the structure of the GR2M model together with relevant equations defining the model structure. (Adapted from Mouelhi *et al.* (2013) and Lespinas *et al.* (2014))

undertaken to determine which parameter sets to retain. First, objective function scores were ranked by their performance, with the top 400 sets from each being retained. Second,

retained simulations were evaluated by their absolute per cent bias (PBIAS) relative to observed flows, with the 200 best performing parameter sets for both logNSE and KGE

retained. The median (henceforth GR2M median) and 95th percentile confidence intervals of GR2M simulated flows, retained from this process, were then determined.

2.3.2 | The ANN Model

ANNs have been widely used for rainfall–runoff modelling (Dawson and Wilby, 1998; Dastorani *et al.*, 2010). A backpropagation ANN was developed here using the neuralnet R package (Fritsch *et al.*, 2019), with different combinations of inputs and neurons tested with two hidden layers. The same calibration and validation periods for individual catchments were employed as those for the GR2M, again using observed data to generate the model. When determining the ANN structure, input data were limited to observed variables that were also available for the full reconstruction period (temperature, precipitation and PET). Lagged variables (e.g. precipitation from previous

months) were also included. The best performing ANN inputs were found to be temperature and precipitation from the current month, plus precipitation lagged by one, two and three months. An example ANN structure which generated the best efficiency scores for one catchment is shown in Figure 5.

Uncertainty in ANN model structure was explored by varying combinations of neurons in one or two hidden layers. Neuron permutations, varying from one to twenty for each hidden layer (giving 420 independent model structures in total), were used to simulate flows for the given calibration period. Each model structure was then independently evaluated using logNSE and KGE and ranked in order of performance. As per the GR2M model, the top 400 ANN model structures according to each objective function were identified and those which subsequently produced the 200 lowest PBIAS scores were retained. The median (henceforth ANN median) and 95th percentile confidence intervals of simulated flows were then obtained.

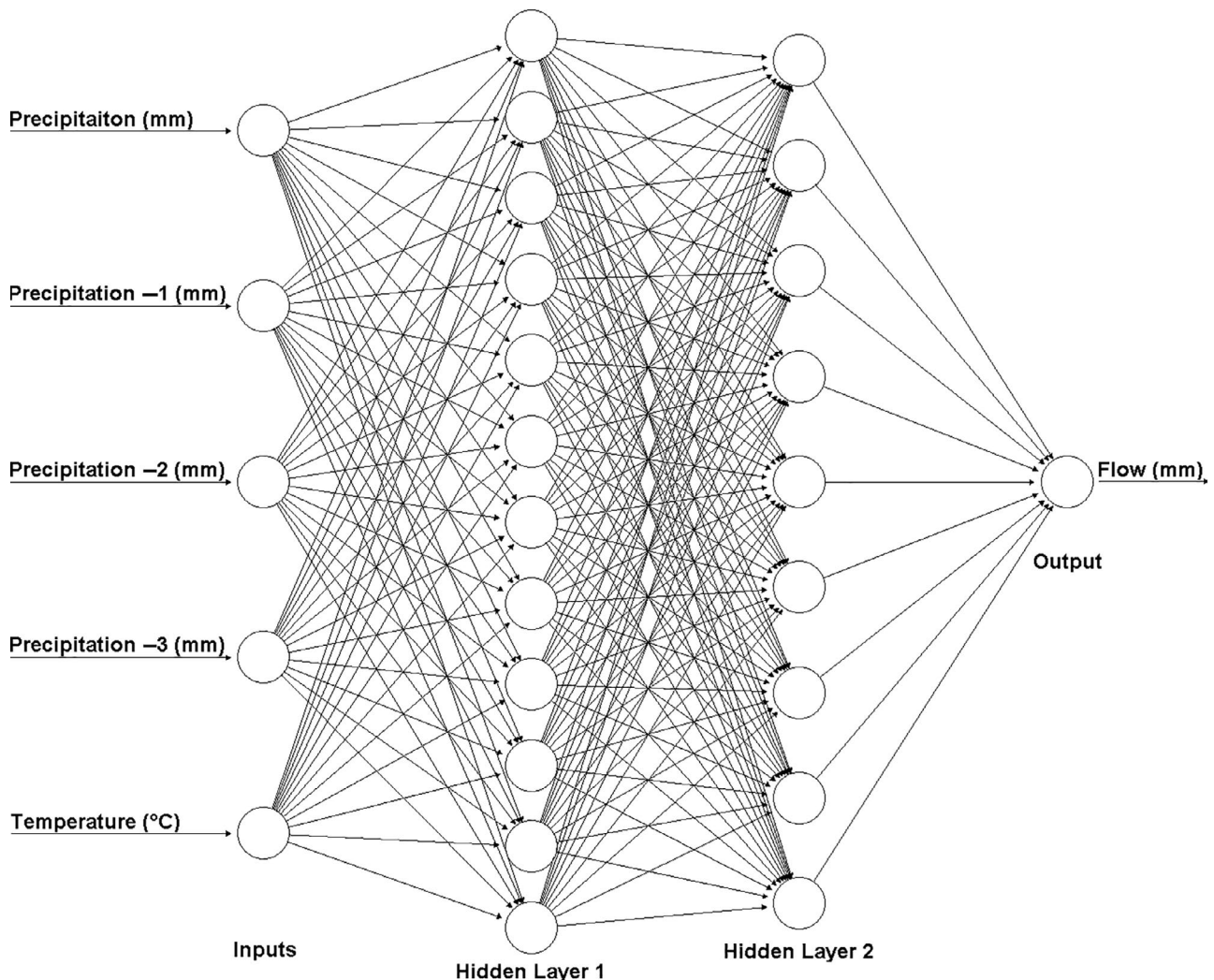


FIGURE 5 Schematic of a typical ANN model structure employed with five inputs, two hidden layers (with 12 and 9 neurons respectively) and monthly flow output. Negative one, two and three values represent the number of lagged months for precipitation

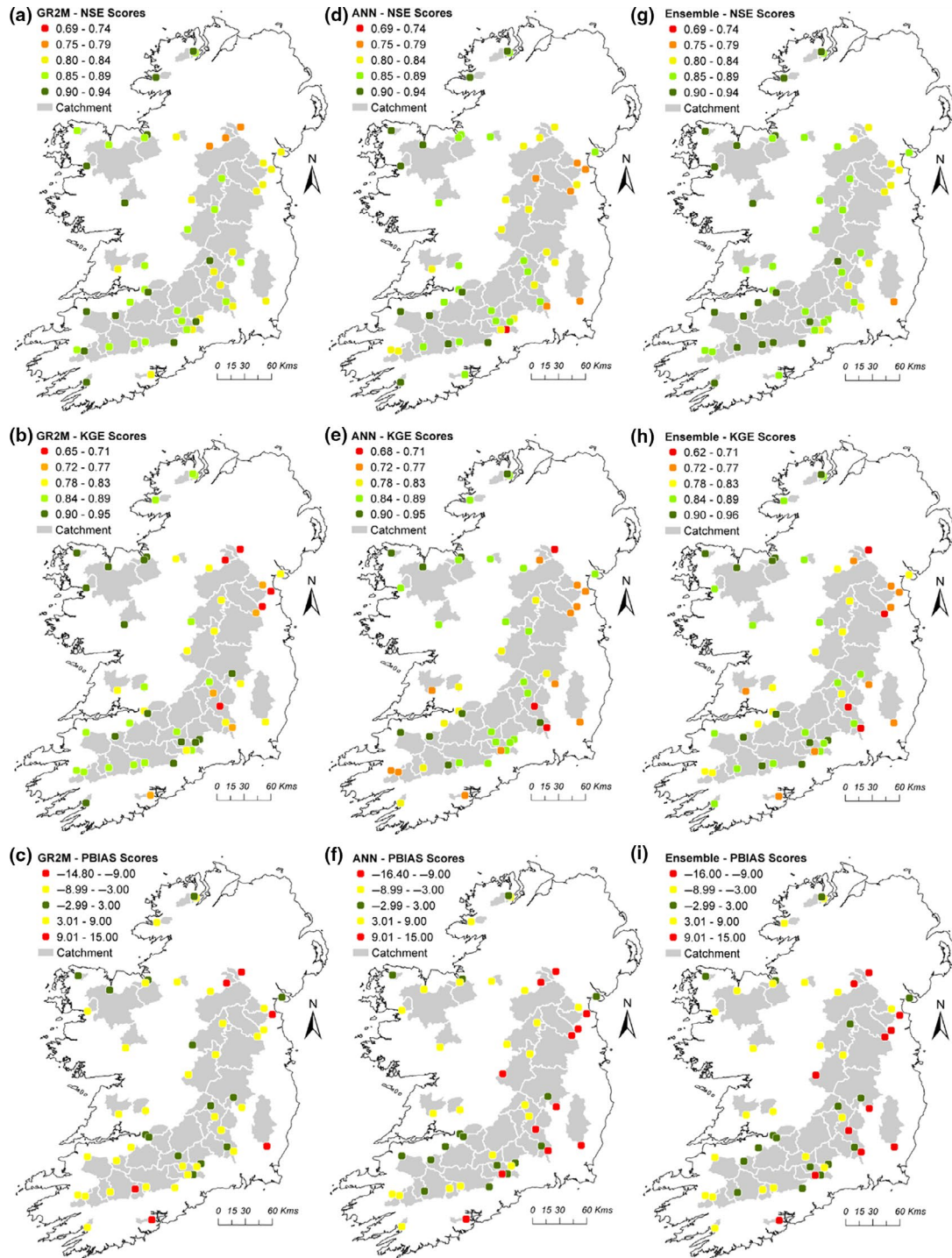


FIGURE 6 Maps of logNSE, KGE and PBIAS scores for GR2M, ANN and Ensemble median simulations for all 51 catchments. Scores are derived from the observed versus modelled flows for the independent validation period (2001–2016) for each catchment

Finally, a mixed ensemble was derived from both GR2M and ANN model structures and parameters by combining the 200 retained simulations from each. The median (henceforth Ensemble median) and 95th percentile confidence intervals of simulated flows were obtained and used to evaluate model reconstructions.

2.3.3 | Validation results

Figure 6 displays the performance of the GR2M, ANN and Ensemble median simulations for all 51 catchments for the 2001–2016 validation period according to logNSE, KGE and PBIAS scores. The ANN and GR2M simulations perform

equally well with average logNSE, KGE and PBIAS scores across all 51 catchments of 0.86, 0.83 and -3.04% for GR2M median and 0.85, 0.83 and -4.97% for ANN median. The combined Ensemble median returned scores of 0.87, 0.83 and -4.38% . Individual catchment results also show similar performance for both model types.

Skill scores for GR2M, ANN and Ensemble median simulations during validation for each catchment are provided in Table 1. Poorest performances are evident for the Nire at Fourmilewater (ID: 16013) which has a logNSE score of 0.69 (ANN median) and the Finn at Anlore (ID: 36015) with a KGE score of 0.65 (GR2M median). PBIAS scores vary between catchments with the largest bias evident for the Blackwater at Faulkland (ID: 3051) (-16.4% ; ANN median) and a minimum of 0% for the Glenamoy at Glenamoy (ID: 33001) (GR2M median). PBIAS values are generally higher for the ANN median.

Observed and simulated monthly flows for the validation period for nine catchments are shown in Figure 7. This subset represents a spread of the best (top row), average (middle row) and worst (bottom row) performing catchments. The proportion of observed variance (R^2) captured by the Ensemble median simulation for each catchment is also provided—varying between 0.88 and 0.93 for the nine sample catchments. The average Ensemble median R^2 value across all 51 catchments for the same validation period is 0.90. ANN and GR2M median simulations show good agreement for the majority of catchments. Whilst observed flows are largely contained within the uncertainty bounds for each of the catchment reconstructions, some discrepancies are apparent in peak values. Arterial drainage works have been identified as a probable cause of this, with previous work showing the tendency for elevated peak flows following drainage (Harrigan *et al.*, 2014). Peak flows also tend to be underestimated for smaller catchments where gridded rainfall may not capture flood generating precipitation adequately.

3 | RECONSTRUCTED FLOWS

3.1 | Assessment of reconstructed flows

Following calibration and validation with observed data, bias corrected Casty data (precipitation/temperature and Oudin PET) were input to the hydrological models to reconstruct monthly river flows back to 1766. The following sub-sections present the resulting annual, seasonal and monthly flow reconstructions across all 51 catchments.

3.1.1 | Annual flow reconstructions

The median of annual reconstructed flows for all 51 catchments from 1766 is shown in Figure 8. GR2M and ANN

median reconstructions show close agreement ($R^2 = 0.97$). In Figure 8 and subsequent plots, observed flows from 1980 onward are displayed as, by this year, observed values are available for over 84% of catchments. Overall, the percentage of median annual observed flow values across all 51 catchments contained within the uncertainty ranges of the median ensemble (henceforth the containment value) is 97%. Observed and Ensemble median simulated series across all catchments show close agreement ($R^2 = 0.81$). Some divergence is evident between modelled and observed flows around 1989 due to differences between Casty and observed precipitation at that time.

3.1.2 | Seasonal and monthly flow reconstructions

Seasonal and monthly flow reconstructions for all 51 catchments are displayed in Figures 9 and 10, respectively, with reconstructions showing strong agreement with observations for 1980–2016 in all seasons. There is some evidence that summer flows are over-estimated in 1989, consistent with annual flows. For all other periods and seasons, observed flows lie within uncertainty estimates (minimum containment value is 89%) and show good agreement with reconstructions (R^2 between Ensemble median values and observations range from a high of 0.9 in summer [JJA] to a low of 0.76 in autumn [SON]). Close agreement is also evident between GR2M and ANN median reconstructions ($R^2 > 0.91$) in all seasons. It is notable from Figure 9 that GR2M reconstructions for spring and summer are slightly higher and autumn values lower than ANN reconstructions.

Monthly reconstructions are displayed in Figure 10 for all 51 catchments. Good agreement is evident between GR2M and ANN median reconstructions ($R^2 > 0.84$ in all months). GR2M median reconstructions are slightly higher than the ANN in April, May, June and July, whilst GR2M output in September, October and November is lower than the ANN equivalent, concurrent with summer and autumn differences between GR2M and ANN values identified above. As expected, performance of monthly simulations is poorer than for seasonal and annual time steps. Monthly observed flows generally lie within uncertainty estimates (mean containment value across all months is 68%) and show satisfactory agreement with observations (R^2 for Ensemble median values vs. observations range between 0.56 in April and 0.91 in July).

3.2 | Comparison with reconstructions from long-term precipitation series

Monthly river flow reconstructions generated with the bias corrected Casty data were evaluated against reconstructions based on monthly precipitation data for stations within the

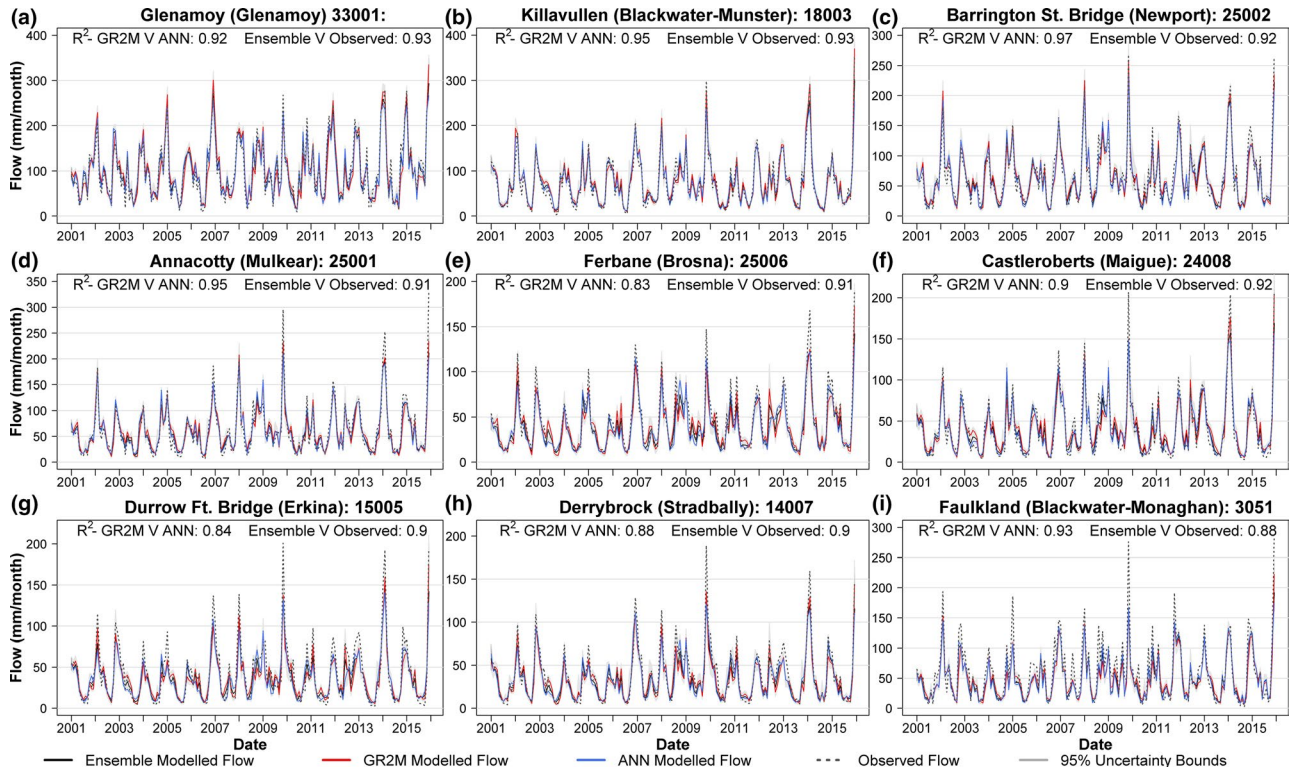


FIGURE 7 Observed and simulated annual mean flows for nine sample catchments representing best (top row), average (middle row) and worst (bottom row) performing models. Plotted are the GR2M (red), ANN (blue) and Ensemble median (black) simulations, together with observed flows (dashed dark-grey). 95% uncertainty range (grey) is derived from the Ensemble median simulations

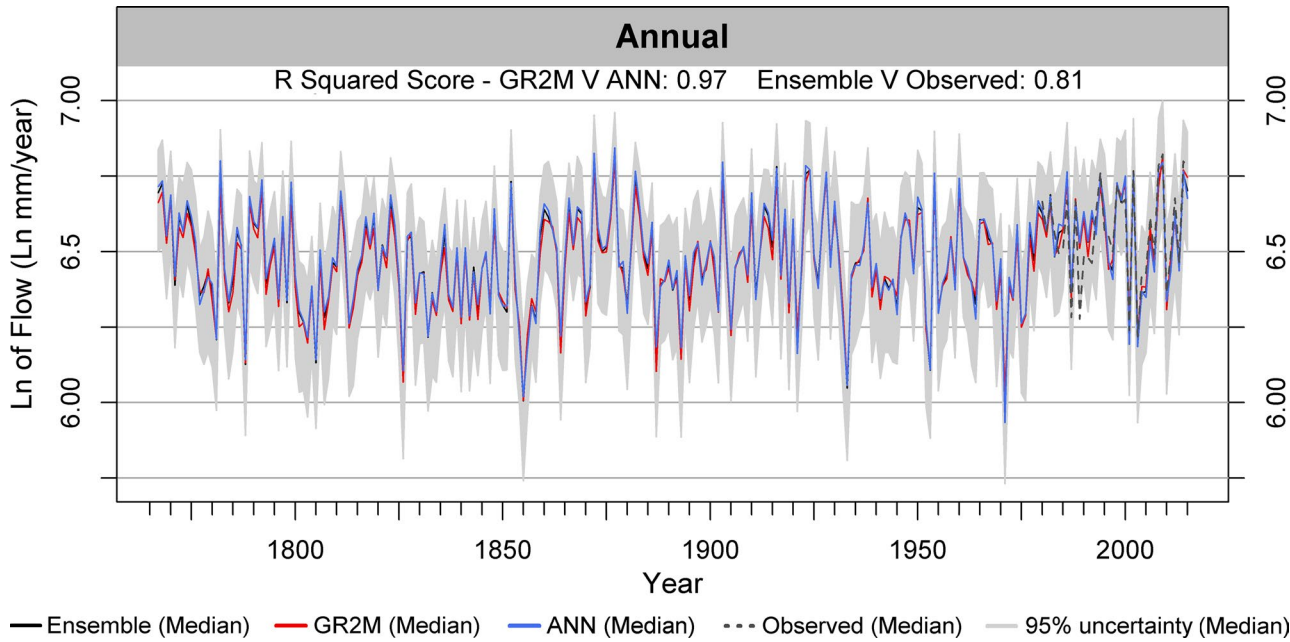


FIGURE 8 Median annual flow values across all 51 catchments for the period 1766–2016 for GR2M (red), ANN (blue) and Ensemble median (black) reconstructions. The median of observed flows across the catchment sample for years 1980–2016 are in dark-grey, whilst 95% uncertainty ranges (grey) are derived from the ensemble simulations

Island of Ireland Precipitation (IIP) network 1850–2010 (Noone *et al.*, 2016). For each catchment, we identified the nearest IIP station (see Figure 1) and then bias corrected data

to catchment average precipitation, as per the Casty data. Bias corrected precipitation, together with bias corrected monthly temperature/PET derived from the Casty data, was

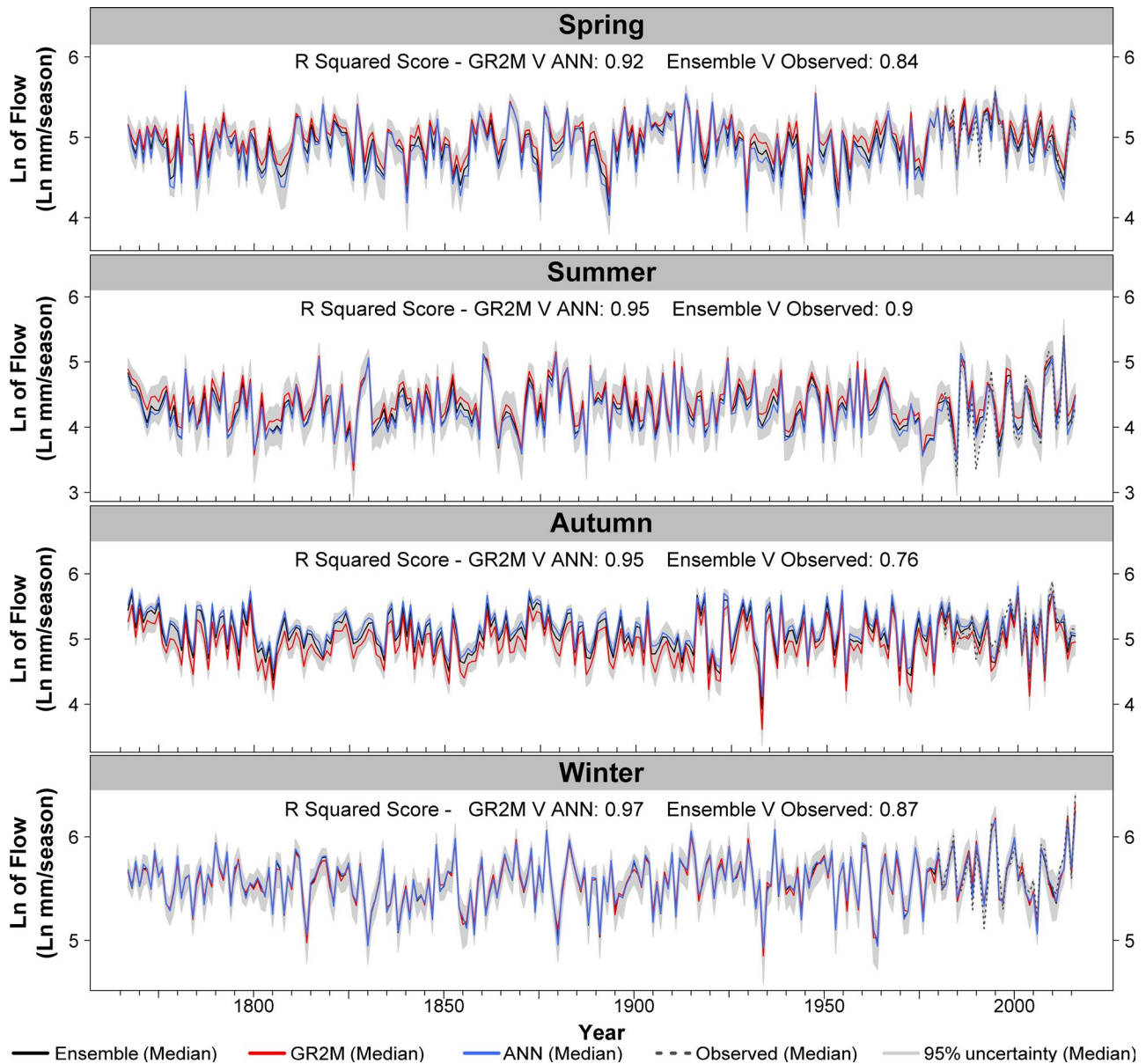


FIGURE 9 As in Figure 8 but for seasonal median flows: Winter [DJF], spring [MAM], summer [JJA], autumn [SON]

used to reconstruct flows back to 1850, using the same methods as described above. Although some of the IIP data are likely contained within the Casty gridded precipitation (so there is a degree of circularity), it was deemed important to compare both data sources, given the different methods used in their construction.

Figure 11 shows the Ensemble median annual mean flow reconstructions from 1850 to 2016 for four exemplar catchments, using Casty precipitation or IIP as input. Strong agreement between the reconstructions is evident despite the different input data with IIP reconstructions largely contained within the uncertainty ranges of the Casty reconstructions. Across the four case study catchments, the R^2 between IIP and Casty reconstructed annual mean flows varies between 0.70 and 0.77. Differences

between flows generated from the two data sources are not unexpected given that IIP data are station based and often located outside catchment boundaries, whereas Casty data are gridded.

3.3 | High- and low-flow assessment

The most notable extreme flow years for seasonal and annual Casty reconstructions were identified (Table 2), with the top five highest and lowest flow year across all catchments displayed for calendar years (1767–2016) as well as winter and summer seasons (1767–2016). The percentage anomaly relative to the mean of the full record is also provided. The most exceptional high-flow years across the

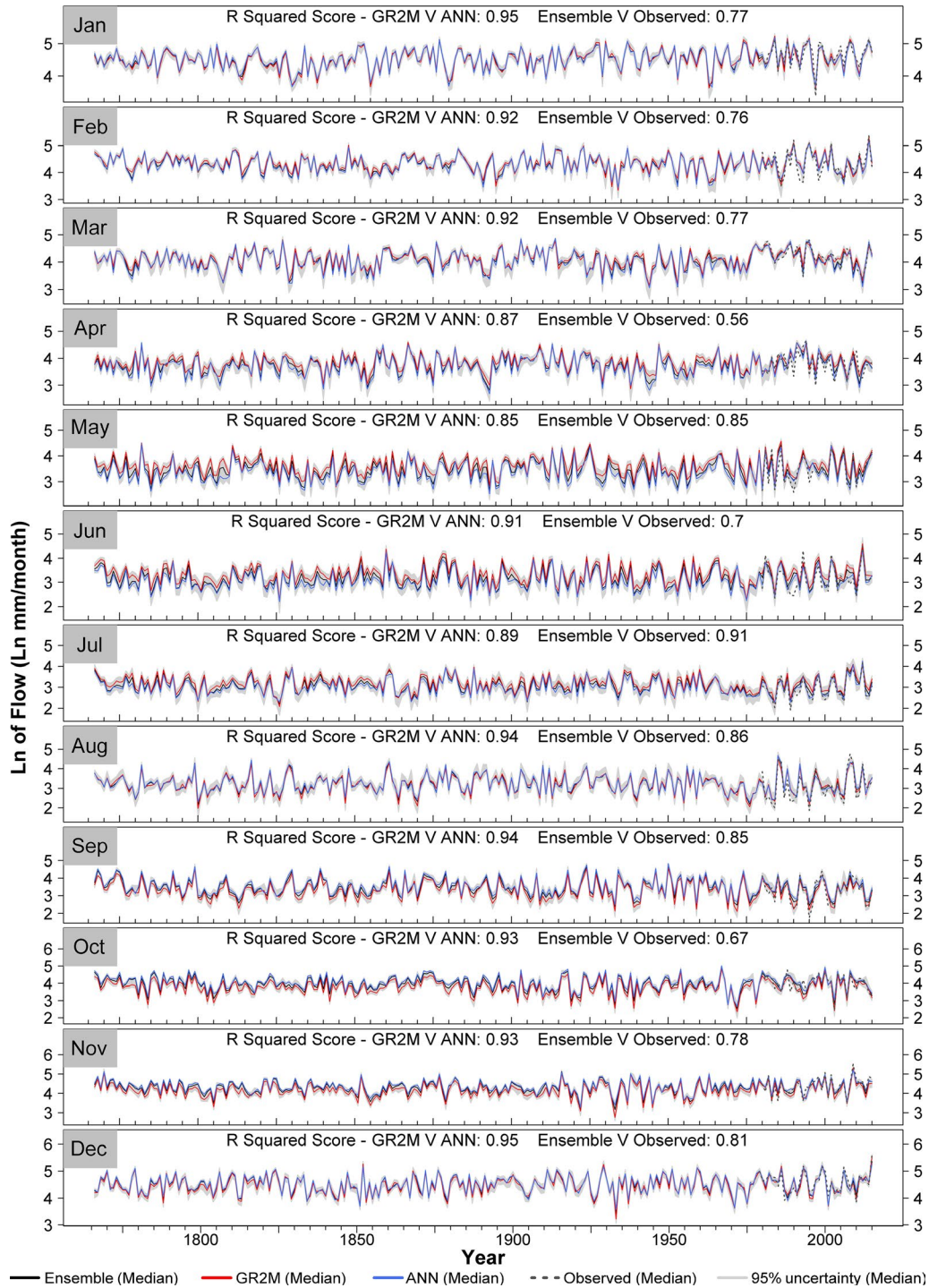


FIGURE 10 As in Figure 8 but for monthly median flows

sample include 1877, 1872 and 1916, whilst the most notable winter seasons include 2015/16, 1994/95 and 2013/14. In terms of exceptional low-flow years, 1855, 1933 and 1971 stand out across the catchments, whereas 1826, 1975 and 1887 dominate the most notable low-flow years for summer. Annual flow anomalies across all 51 catchments range from 150% to 58% of the long-term mean for all catchments, whilst seasonally winter and summer extreme

anomalies range from 173% to 37% of the respective long-term seasonal mean values. Our extreme years and seasons show considerable agreement with a similar evaluation of reconstructed river flows (1865–2002) in the United Kingdom (Jones *et al.*, 2006), with the previously identified exceptional high- and low-flow seasons and years (1865–2002) all found at least once in the top five equivalent events for multiple catchments in that series.

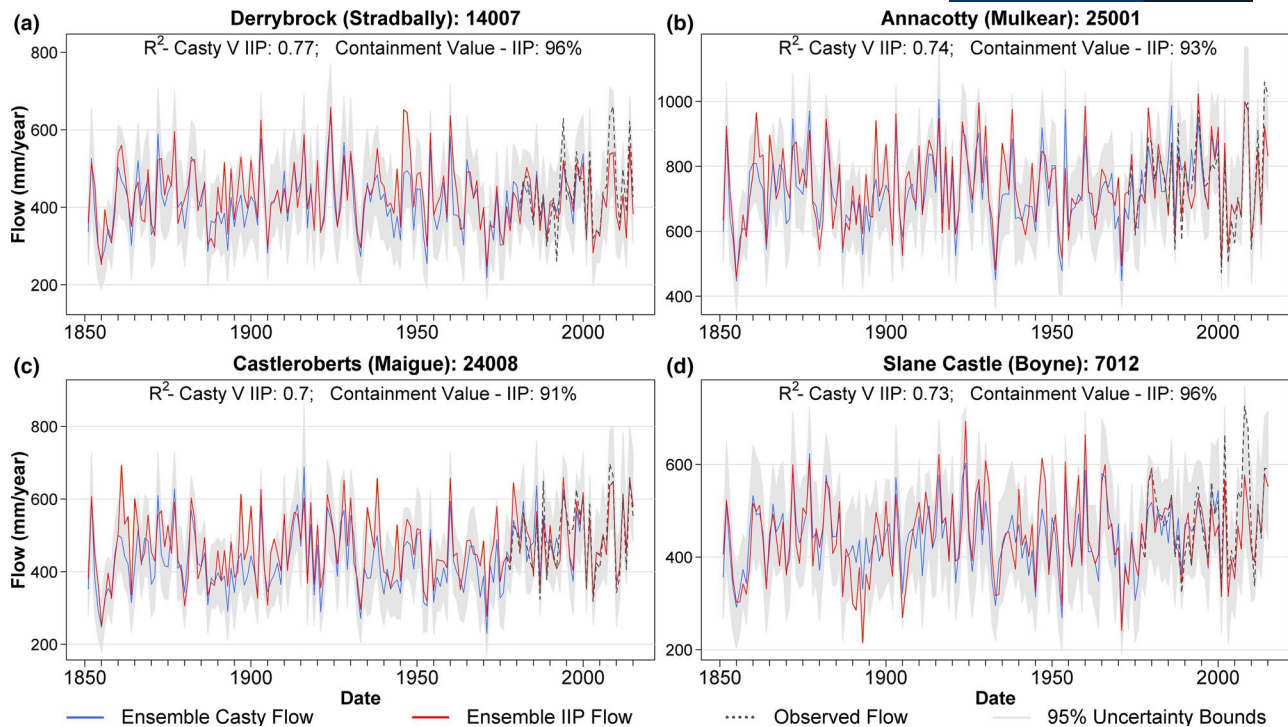


FIGURE 11 Reconstructed annual mean flow values for four sample catchments. Ensemble median simulations generated using Casty precipitation data (blue), and Island of Ireland Precipitation (IIP) data (red), together with observed flows (dashed dark-grey) are displayed for each catchment

4 | DATA SET ACCESS, USES AND LIMITATIONS

The derived monthly flow reconstructions (December 1766 to November 2016 inclusive) for the 51 catchments are freely available for download from the PANGAEA data centre (<https://doi.org/10.1594/PANGAEA.914306>). Data are presented as five individual tab-delimited text files (ASCII), representing reconstructions for each catchment from the GR2M, ANN and Ensemble median simulations, along with 2.5% and 97.5% quantiles derived from the Ensemble simulation. Also included is a table providing the geographical co-ordinates of all 51 flow stations.

4.1 | Potential uses

The reconstructed flow series provide a resource for assessing the impacts of extreme meteorological events, such as drought, on river flows across Ireland, extending the work of Noone *et al.* (2017) and Noone and Murphy (2020). Our reconstructions could also inform spatio-temporal assessments of variability plus support detection of multi-centennial changes in river flows (e.g. Wilby, 2006). Furthermore, the multi-centennial time scale of our reconstructions offers the potential to examine how modes of ocean and climate variability influence river flows over extended periods. For example, it is known that Atlantic multidecadal variability

exerts an important control on Ireland's climate (McCarthy *et al.*, 2015), but its impact on river flows is less clear. Our long-term data set offers the means to explore any potential control, including its stationarity. In turn, this could help facilitate improved seasonal forecasting (e.g. Wedgbrow *et al.*, 2002).

This work represents the first reconstruction of monthly flows for a large number of Irish catchments using long-term reanalysis data and observations. Given the uncertainties involved, this data set should be treated as a benchmark and evaluated and improved by future products. The approach to flow reconstruction adopted here is easily transferable to other catchments in Europe (i.e. the domain of Casty data). By taking advantage of observed runoff data, available from the Global Runoff Data Centre (https://www.bafg.de/GRDC/EN/Home/homepage_node.html), it would be possible to generate similar archives of monthly flow reconstructions for the entire continent.

4.2 | Limitations

There are several recognized limitations to reconstructed river flows. First, arterial drainage has had a pervasive impact on Irish rivers. Catchments in this data set that have been drained tend to have higher peak flows during winter months than captured by the reconstructions. This is consistent with the findings of Harrigan *et al.* (2014) for the

TABLE 2 Years with the five highest and lowest annual (calendar), winter [DJF] (year given for January) and summer [JJA] flows for the period of reconstructions 1767–2016 across all 51 catchments

Station (ID)	Top 5 High Flow Years										Top 5 Low Flow Years									
	Annual					Winter (December to February)					Annual					Summer (June to August)				
	1	2	3	4	5	1	2	3	4	5	1	2	3	4	5	1	2	3	4	5
3051	2015	2002	1877	1954	2009	2016	1995	1994	2014	1937	1933	1826	1855	1953	1971	1826	1975	1826	1983	1870
	150%	145%	144%	138%	137%	193%	172%	166%	161%	159%	59%	62%	62%	65%	59%	33%	33%	35%	37%	39%
6013	1877	1966	2002	1782	1872	2016	1877	1994	1883	1915	1953	1971	1933	1975	1855	1975	1826	1887	1995	1870
	154%	150%	146%	141%	138%	170%	158%	154%	151%	151%	53%	57%	62%	63%	64%	30%	32%	38%	39%	42%
6014	2002	1877	1966	1782	1872	2016	1877	1883	1994	1915	1953	1971	1855	1933	1826	1975	1826	1995	1887	1870
	160%	152%	146%	138%	136%	174%	153%	151%	151%	148%	55%	57%	63%	63%	64%	35%	36%	43%	44%	47%
6030	2002	1877	2015	1782	1872	2016	2013	1994	1877	1990	1933	1826	1953	1887	1911	1995	1800	1826	1949	1983
	173%	142%	141%	137%	137%	201%	173%	163%	161%	158%	61%	64%	65%	67%	69%	24%	28%	28%	31%	31%
7009	1877	1924	1782	1872	1960	2016	2014	1877	1995	1994	1971	1953	1855	1933	1826	1975	1826	1995	1887	1870
	142%	137%	135%	135%	133%	173%	151%	147%	146%	145%	57%	61%	66%	67%	68%	36%	38%	42%	44%	45%
7012	1877	1924	1782	1872	1965	2016	1877	2014	1937	1995	1971	1953	1855	1933	1826	1975	1826	1887	1995	1870
	152%	141%	140%	138%	138%	180%	164%	163%	161%	158%	56%	60%	64%	64%	65%	32%	34%	41%	41%	43%
12001	1960	1930	1872	1924	2009	2016	1877	1995	1994	1930	1971	1953	1905	1788	1855	1975	1826	1995	1887	1984
	158%	150%	148%	146%	145%	192%	177%	177%	170%	169%	55%	62%	64%	66%	66%	37%	38%	42%	44%	44%
14007	1924	1877	1872	1960	1903	1915	2014	1995	2016	1883	1971	1953	1855	1788	1933	1975	1826	1995	1887	1870
	155%	145%	144%	143%	141%	171%	169%	168%	165%	160%	53%	62%	64%	66%	67%	38%	40%	43%	45%	47%
14019	1924	1960	2008	1877	1872	2014	2016	1995	1915	1994	1971	1953	1855	1905	1933	1975	1826	1995	1887	1870
	149%	144%	142%	141%	140%	175%	173%	170%	165%	162%	57%	65%	67%	69%	70%	43%	46%	47%	52%	53%
15001	1877	1872	1924	1928	1903	1930	1995	1877	1915	1937	1971	1953	1855	1788	1826	1826	1975	1984	1870	1887
	160%	155%	153%	152%	151%	193%	192%	190%	190%	184%	51%	54%	59%	60%	64%	32%	33%	35%	36%	36%
15003	1924	2009	1872	1877	1916	2014	1995	1915	2016	1937	1971	1953	1788	1855	1933	1826	1975	1870	1887	1995
	150%	148%	147%	147%	147%	179%	177%	173%	164%	162%	50%	57%	61%	64%	64%	24%	24%	27%	28%	28%
15005	1877	1872	1903	1916	1924	1930	1915	1937	2016	1877	1953	1971	1855	1788	1933	1826	1984	1887	1975	1870
	152%	145%	142%	142%	141%	180%	175%	174%	174%	171%	57%	61%	62%	67%	67%	39%	40%	43%	43%	46%
15006	1877	1872	1924	1928	1903	1915	1995	2016	1930	1877	1971	1953	1855	1788	1933	1826	1975	1984	1887	1870
	152%	146%	145%	144%	143%	178%	178%	176%	174%	172%	57%	60%	65%	66%	68%	39%	41%	42%	44%	45%
15007	1872	1877	1954	1903	1924	2016	1995	1937	2014	1915	1971	1953	1855	1933	1826	1826	1984	1800	1870	1975
	142%	142%	141%	140%	140%	177%	173%	172%	172%	168%	57%	58%	59%	61%	65%	37%	39%	43%	43%	44%

(Continues)

TABLE 2 (Continued)

Station (ID)	Top 5 High Flow Years										Top 5 Low Flow Years									
	Annual					Winter (December to February)					Annual					Summer (June to August)				
	1	2	3	4	5	1	2	3	4	5	1	2	3	4	5	1	2	3	4	5
16008	1877	1872	1903	1923	1954	2016	1877	1995	1877	2014	1953	1855	1971	1933	1826	1984	1826	1870	1887	1975
	144%	143%	140%	139%	139%	180%	168%	167%	161%	161%	58%	59%	60%	65%	67%	38%	40%	44%	47%	48%
16009	1877	2009	1872	1903	1916	2016	1995	1937	1915	2014	1855	1971	1953	1933	1826	1826	1984	1870	1975	1887
	147%	144%	143%	143%	142%	189%	184%	175%	172%	170%	60%	61%	62%	66%	68%	45%	45%	51%	51%	52%
16010	2009	1877	1928	1903	1916	1930	2016	1915	1995	1994	1855	1971	1788	1953	1893	1826	1984	1870	1975	1887
	153%	152%	152%	148%	145%	189%	189%	187%	181%	179%	61%	64%	66%	67%	68%	44%	47%	50%	50%	51%
16011	1877	1903	2009	1928	1872	2016	1995	1915	1994	1937	1971	1855	1953	1933	1788	1826	1984	1975	1870	1887
	143%	141%	141%	140%	139%	185%	181%	172%	168%	167%	64%	65%	67%	69%	71%	51%	52%	56%	57%	57%
16012	1928	2009	1877	1903	1960	1995	2016	1915	1930	1925	1855	1971	1826	1933	1887	1826	1984	1887	1975	1870
	143%	141%	140%	138%	138%	188%	187%	170%	169%	168%	65%	66%	70%	71%	72%	47%	50%	53%	53%	56%
16013	1928	2009	1938	1903	1960	1930	1995	2016	1915	1994	1971	1933	1855	1788	1887	1826	1984	1870	1800	1887
	152%	148%	147%	145%	145%	196%	185%	185%	184%	182%	63%	66%	67%	68%	69%	38%	40%	43%	45%	46%
18002	1916	2009	1877	2000	1872	2016	2014	1995	1915	1994	1971	1855	1933	1788	1826	1826	1984	1887	1975	1870
	151%	142%	140%	139%	135%	209%	188%	184%	163%	162%	56%	61%	69%	70%	71%	44%	45%	50%	50%	51%
18003	1916	2009	2000	1877	1994	2016	2014	1995	1915	1994	1971	1855	1788	1826	1921	1984	1826	1887	1975	1870
	150%	143%	141%	138%	138%	215%	199%	189%	170%	167%	54%	61%	67%	69%	69%	38%	39%	45%	45%	47%
18006	1916	2009	2000	1994	2002	2016	2014	1995	1915	1994	1971	1855	1826	1933	1788	1984	1826	1887	1975	1800
	145%	144%	141%	140%	139%	217%	205%	191%	177%	173%	58%	62%	68%	69%	70%	42%	43%	47%	47%	48%
18050	2009	2002	2008	2000	1872	2016	2014	1995	1915	1994	1971	1855	1788	1826	1921	1826	1984	1800	1975	1944
	142%	139%	139%	138%	137%	203%	200%	185%	179%	172%	58%	62%	69%	69%	70%	40%	41%	42%	43%	44%
19001	2009	1982	1928	2000	1872	2016	1995	2014	1915	1994	1971	1855	1788	1933	1854	1826	1826	1887	1870	1975
	159%	153%	151%	150%	149%	233%	211%	203%	188%	187%	48%	57%	63%	63%	65%	38%	38%	40%	41%	42%
21002	2009	1982	1872	1914	1916	2016	1995	1915	2014	1994	1971	1855	1955	1788	1887	1800	1976	1955	1975	1864
	142%	136%	134%	134%	134%	203%	178%	168%	167%	164%	60%	70%	71%	72%	72%	33%	39%	42%	42%	46%
22006	1916	1872	1982	2000	2002	2016	2014	1995	1915	1994	1971	1855	1933	1788	1921	1800	1976	1975	1984	2006
	137%	135%	135%	135%	135%	199%	187%	183%	175%	169%	62%	66%	69%	71%	72%	45%	47%	51%	54%	54%
22035	2000	1872	1916	1982	1994	2016	2014	1995	1915	1994	1971	1855	1933	1788	1921	1800	1975	1984	1826	1976
	137%	136%	135%	135%	131%	202%	182%	176%	168%	163%	62%	67%	72%	73%	73%	51%	53%	56%	57%	58%
23002	2008	2015	1986	1916	1872	2014	2016	1995	1915	1994	1971	1855	1788	1826	1921	1800	1984	1826	1975	1976

TABLE 2 (Continued)

Station (ID)	Top 5 High Flow Years															Top 5 Low Flow Years															
	Annual					Winter (December to February)					Annual					Summer (June to August)															
	1	2	3	4	5	1	2	3	4	5	1	2	3	4	5	1	2	3	4	5											
24008	157%	145%	141%	140%	138%	215%	205%	195%	187%	177%	56%	1971	1855	1933	1788	1826	68%	1826	1975	1887	1870	39%	1826	1975	1887	1870	39%	1826	1975	1887	1870
24030	160%	154%	151%	148%	148%	228%	224%	212%	180%	176%	53%	1971	1855	1933	1788	1826	66%	1826	1975	1887	1870	42%	1826	1975	1887	1870	43%	1826	1975	1887	1870
25001	156%	153%	151%	151%	150%	228%	226%	213%	189%	188%	51%	1971	1855	1933	1788	1826	64%	1826	1975	1887	1870	42%	1826	1975	1887	1870	44%	1826	1975	1887	1870
25002	141%	140%	138%	137%	136%	185%	180%	177%	163%	161%	63%	1971	1855	1933	1788	1826	67%	1826	1975	1887	1870	47%	1826	1975	1887	1870	48%	1826	1975	1887	1870
25006	144%	138%	137%	135%	134%	186%	167%	164%	163%	162%	62%	1971	1855	1933	1788	1826	67%	1826	1975	1887	1870	38%	1826	1975	1887	1870	39%	1826	1975	1887	1870
25030	157%	145%	144%	140%	139%	219%	210%	201%	177%	162%	59%	1971	1855	1933	1788	1826	63%	1826	1975	1887	1870	41%	1826	1975	1887	1870	42%	1826	1975	1887	1870
25034	145%	140%	140%	140%	139%	190%	165%	158%	154%	152%	54%	1971	1855	1933	1788	1826	67%	1826	1975	1887	1870	35%	1826	1975	1887	1870	40%	1826	1975	1887	1870
26021	141%	139%	139%	138%	135%	196%	163%	159%	153%	148%	58%	1971	1855	1933	1788	1826	65%	1826	1975	1887	1870	42%	1826	1975	1887	1870	44%	1826	1975	1887	1870
26029	144%	138%	134%	133%	131%	183%	173%	164%	154%	152%	64%	1971	1855	1933	1788	1826	65%	1826	1975	1887	1870	34%	1826	1975	1887	1870	35%	1826	1975	1887	1870
26058	2002	1877	1924	2015	2014	2016	2014	1937	2007	1995	1971	1855	1933	1826	1805	1805	1826	1975	1887	1870	34%	1826	1975	1887	1870	35%	1826	1975	1887	1870	
27002	143%	140%	139%	138%	136%	195%	166%	153%	144%	143%	58%	1971	1855	1933	1788	1826	62%	1826	1975	1887	1870	42%	1826	1975	1887	1870	43%	1826	1975	1887	1870
30007	160%	150%	148%	139%	139%	221%	199%	192%	168%	165%	57%	1971	1855	1933	1788	1826	63%	1826	1975	1887	1870	39%	1826	1975	1887	1870	40%	1826	1975	1887	1870
32012	154%	148%	140%	140%	138%	197%	169%	165%	164%	153%	58%	1971	1855	1933	1788	1826	62%	1826	1975	1887	1870	39%	1826	1975	1887	1870	46%	1826	1975	1887	1870
33001	147%	142%	137%	130%	130%	180%	178%	165%	154%	151%	64%	1971	1855	1933	1788	1826	65%	1826	1975	1887	1870	48%	1826	1975	1887	1870	50%	1826	1975	1887	1870
	149%	141%	131%	128%	127%	185%	177%	166%	151%	147%	65%	1971	1855	1933	1788	1826	66%	1826	1975	1887	1870	34%	1826	1975	1887	1870	35%	1826	1975	1887	1870

(Continues)

TABLE 2 (Continued)

Station (ID)	Top 5 High Flow Years															Top 5 Low Flow Years														
	Annual					Winter (December to February)					Annual					Summer (June to August)														
	1	2	3	4	5	1	2	3	4	5	1	2	3	4	5	1	2	3	4	5										
34001	2015	1986	1950	2008	1877	2016	1995	1937	1950	1855	1933	1805	1826	2003	1826	1826	1984	1975	1800	1887										
	143%	141%	138%	138%	136%	185%	159%	156%	149%	60%	63%	68%	69%	70%	69%	41%	45%	49%	51%	51%										
35002	2015	2008	2002	1986	1949	2016	1995	2015	1994	1855	1933	1805	1826	1921	1826	1800	1826	1995	1983	1984										
	150%	146%	145%	136%	132%	187%	167%	156%	154%	63%	64%	70%	70%	75%	70%	40%	43%	43%	47%	47%										
35005	2015	2002	2008	1986	1950	2016	1937	1995	2015	1855	1933	1805	1826	1921	1826	1826	1984	1800	1995	1975										
	149%	144%	143%	140%	137%	190%	164%	163%	153%	62%	63%	66%	66%	69%	66%	38%	42%	46%	46%	47%										
36015	1877	1954	1998	2002	1782	1995	1937	1994	1877	1933	1826	1855	1971	1921	1975	1826	1983	1983	1800	1870										
	146%	145%	142%	141%	138%	169%	158%	157%	153%	57%	60%	60%	63%	65%	27%	27%	28%	31%	32%	33%										
36019	2002	1877	2015	1954	1998	2016	1937	1995	2007	1971	1933	1826	1855	1953	1826	1826	1975	1984	1870	1887										
	155%	146%	141%	139%	138%	179%	158%	154%	153%	58%	59%	61%	61%	65%	24%	24%	24%	31%	33%	34%										
38001	2015	1949	2011	1992	1986	1995	2016	1994	2015	1933	1855	1805	1826	1864	1800	1800	1821	1995	1983	1984										
	145%	139%	136%	132%	131%	183%	165%	162%	160%	63%	64%	69%	70%	75%	29%	29%	37%	37%	40%	40%										
39006	2015	1990	1949	1992	1999	2016	2015	2014	1994	1855	1933	1826	1805	1911	1800	1800	1995	1983	1824	1984										
	167%	141%	138%	138%	137%	189%	179%	176%	173%	60%	63%	66%	67%	70%	34%	34%	36%	36%	37%	38%										
39009	2015	1949	1990	1992	1999	2016	2015	2014	1994	1855	1933	1805	1826	1911	1800	1800	1983	1995	1984	1824										
	165%	140%	140%	138%	138%	186%	177%	175%	173%	60%	62%	67%	67%	71%	32%	32%	34%	34%	35%	36%										

Note: The percentage anomaly relative to the long-term mean (1767–2016) is provided in each case. Values highlighted in progressively darker blue represent the top three occurring high flow events, whilst those in red represent the top three occurring low-flow events.

Boyne catchment. Hence, our reconstructions may be useful for quantifying the impact of arterial drainage on flow response. Moreover, we note that there is limited knowledge about how arterial drainage affects low-flow and drought responses—again, our reconstructions may provide a useful point of reference.

Changes in land use can have considerable impacts on flows over time (Yan *et al.*, 2013). Lack of metadata on historical land-use change hinders the quantification of such impacts. Moreover, Slater *et al.* (2019) highlight that rivers are treated as conduits of fixed conveyance by models even though changes in channel geometry and structure are known to occur in response to periods of hydro-climatic variability. Here, we assume that land-use and channel geomorphology remain static over the period of reconstruction; a common assumption attached to long-term flow reconstructions. Jones (1984) asserts that such assumptions can be justified. Water resource infrastructure designs are based on flows relating to current land use as opposed to historical conditions, suggesting that catchment response tuned to present conditions are a useful resource.

Second, potential biases or inaccuracies in precipitation data could propagate into the reconstructed flow series. The gridded Casty data set employed in this study was generated using both reanalysis and observed precipitation values, with principle component regression to interpolate across space. Interpolation of station data is more uncertain before the 1900s as the number of stations decreases rapidly prior to this time. Casty *et al.* (2005) highlight that European wide precipitation patterns in the early part of their series should be treated with caution, especially before 1800 when station numbers are low. For Ireland, we believe that data prior to 1850 should be treated with caution due to the sparseness of observed precipitation records on the island. A further source of uncertainty relates to the quality of early precipitation observations. Murphy *et al.* (2019) show that pre-1870 winter precipitation observations in the United Kingdom were likely affected by under-catch of snowfall due to gauge design and observer practice. It is likely that early Irish precipitation totals are affected by the same biases during winter months (Murphy *et al.*, 2020).

Third, the sensitivity of hydrological model parameters to prevailing climatic conditions during the calibration period can result in uncertainties when models are used to simulate conditions different to those used for training. Broderick *et al.* (2016) showed that changes in climatic conditions can affect model performance depending on catchment, model type and assessment criteria. A shift from relatively wet to dry conditions resulted in poorer results. Future work should assess the robustness of monthly reconstructions to the wetness or dryness of periods used for training.

5 | SUMMARY

This paper presents a data set of monthly river flow reconstructions back to 1766 for 51 Irish catchments. Gridded reconstructions of monthly precipitation and temperature, bias corrected to observed catchment data sets, are used with derived PET to force a conceptual hydrological model and an Artificial Neural Network to generate monthly flows spanning more than 250 years. Reconstructed flows are subject to uncertainties associated with hydrological response to arterial drainage and land-use change, together with potential biases in early precipitation observations and non-stationary hydrological model parameters. With these caveats in mind, the data set is suitable for examining hydrological responses to arterial drainage, tracking hydrological variability and change, or testing the robustness of water plans and/or contextualizing modern hydrological droughts.

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OPEN PRACTICES

This article has earned an Open Data badge for making publicly available the digitally-shareable data necessary to reproduce the reported results. The data is available at <https://doi.org/10.1594/PANGAEA.914306> Learn more about the Open Practices badges from the Center for OpenScience: <https://osf.io/tvyxz/wiki>.

ORCID

Paul O'Connor  <https://orcid.org/0000-0002-7755-0831>
 Conor Murphy  <https://orcid.org/0000-0003-4891-2650>
 Tom Matthews  <https://orcid.org/0000-0001-6295-1870>
 Robert L. Wilby  <https://orcid.org/0000-0002-4662-9344>

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